

Attenuation correction for a high-resolution polarimetric X-band weather radar

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Abstract. In 2007, IRCTR (Delft University of Technology) installed a new polarimetric X-band LFM CW radar (IDRA) at the meteorological observation site of Cabauw, The Netherlands. It provides plan position indicators (PPI) at a fixed elevation with a high range resolution of either 3 m or 30 m at a maximum observation range of 1.5 km and 15 km, respectively. IDRA aims to monitor precipitation events for the long-term analysis of the hydrological cycle. Due to the specifications of IDRA, the spatial and temporal variability of a large range of rainfall intensities (from drizzle to heavy convective rain) can be studied.

Even though the usual observation range of IDRA is limited to 15 km, attenuation due to precipitation can be large enough to seriously affect the measurements. In this contribution we evaluate the application of a combined method to correct for the specific and the differential attenuation, and in the same vein estimate the parameters of the raindrop-size distribution. The estimated attenuations are compared to a phase constraint attenuation correction method.

1 Introduction

Weather radars are a widely used measurement device to observe the spatial distribution and the temporal evolution of precipitation. In Europe over 150 operational C-band weather radars are employed by the meteorological services¹. Nowadays, the operational weather radar networks are complemented by compact X-band systems as gap fillers, McLaughlin et al. (2009). However, the attenuation of microwaves by precipitation is stronger at X-band than at C-

band. Radar measurements must be corrected for attenuation before they are used for a quantitative determination of precipitation parameters. A first formulation of attenuation correction for weather radar measurements was provided by Hitschfeld and Bordan (1954). However, without constraining the path-integrated attenuation (PIA) by independent measurements, e.g. rain gauges or microwave links, the attenuation correction is unstable. Polarimetric weather radars operating at linear horizontal/vertical polarisation basis have the capability to measure the differential propagation phase that is a forward scattering effect like attenuation. Scattering computations show that the differential propagation phase is linearly related to the attenuation, Holt (1988), and can those be used to provide an estimate of the PIA. However, the relation between the differential propagation phase and the attenuation depends on the microphysics of the precipitation, Bringi et al. (1990). Testud et al. (2000) introduced an algorithm that uses the differential propagation phase to estimate the PIA that is subsequently distributed among the range bins weighted by the reflectivity. Additionally, self-consistency among the weather radar observables involved in the attenuation correction may be imposed in order to account for the variability of the relation among the differential propagation phase and the attenuation, Bringi et al. (2001), Gorgucci and Baldini (2007).

Attenuation correction is also a prerequisite for IDRA, the polarimetric, horizontally scanning X-band weather radar of IRCTR installed at the meteorological observatory at Cabauw², The Netherlands. One purpose of IDRA is the long-term monitoring of microphysics of precipitation. The method proposed in Chandra et al. (1990) is able to jointly perform the attenuation correction and the raindrop-size distribution retrieval for rain. It was successfully applied to C-band weather radar data. The purpose of this contribution is the evaluation of this method for X-band weather radar data



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¹OPERA – Operational programme for the exchange of weather radar information, <http://www.knmi.nl/opera>

²<http://www.cesar-database.nl>

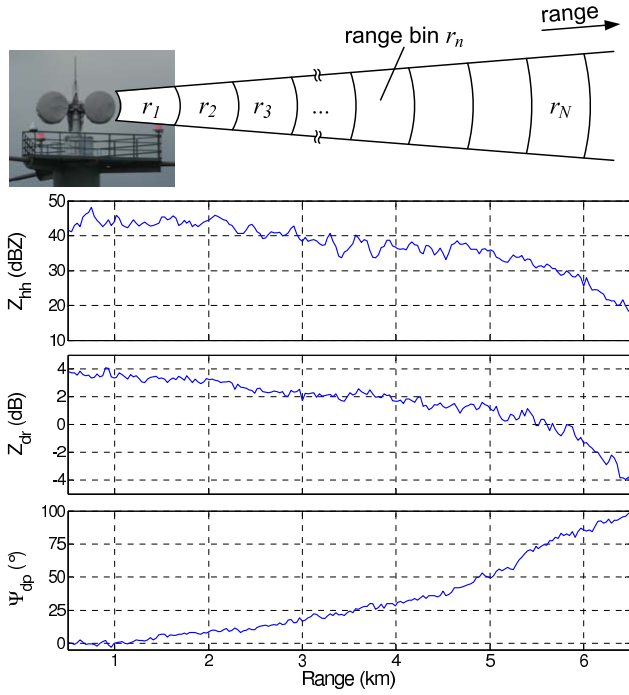


Fig. 1. The radar geometry is shown in the upper part of this figure. The lower part shows rayplots of an IDRA measurement (26 May 2009, 02:00 UTC) with heavy precipitation leading to a very strong specific and differential attenuation that deteriorate the measured reflectivity $Z_{hh,att}$ and the differential reflectivity $Z_{dr,att}$ (upper two ray plots). The measured differential propagation phase shows the expected range-cumulative behaviour (ray plot at the bottom).

by applying it to IDRA data. Therefore, the attenuations derived by this method are compared to the phase constraint attenuation correction method with imposed self-consistency, Gorgucci and Baldini (2007).

2 Basics

The natural variability of rainfall can be described by the three-parameter raindrop-size distribution (RDSD) that normalised with respect to the liquid water content can be written as, Bringi and Chandrasekar (2001),

$$N(D) = N_w f(\mu) \left(\frac{D}{D_0} \right)^\mu e^{-(3.67+\mu)\frac{D}{D_0}} \quad (1)$$

with

$$f(\mu) = \frac{6}{3.67^4} \frac{(3.67+\mu)^{\mu+4}}{\Gamma(\mu+4)} \quad (2)$$

where D is the equivolumetric raindrop diameter, D_0 is the median volume diameter, N_w is the concentration, and μ is the shape parameter.

The measurements of an homogeneous ensemble of raindrops by polarimetric weather radars can be expressed in

terms of the RDSD, and the scattering amplitudes of the raindrops as follows:

$$Z_{hh} = 10 \log_{10} \left(\frac{\lambda^4}{\pi^5 |K|^2} \int_D \sigma_{hh}(D) N(D) dD \right) \quad (3)$$

$$Z_{dr} = 10 \log_{10} \left(\frac{\int_D \sigma_{hh}(D) \left(\frac{D}{D_0} \right)^\mu e^{-(3.67+\mu)\frac{D}{D_0}} dD}{\int_D \sigma_{vv}(D) \left(\frac{D}{D_0} \right)^\mu e^{-(3.67+\mu)\frac{D}{D_0}} dD} \right) \quad (4)$$

$$K_{dp}^{1-way} = \frac{180}{\pi} \lambda \int_D \Re[f_{hh}(D) - f_{vv}(D)] N(D) dD. \quad (5)$$

The intrinsic weather radar observables in Eqs. (3) to (5) represent the co-polarised reflectivity at horizontal polarisation Z_{hh} (dBZ), the differential reflectivity Z_{dr} (dB), and the specific differential phase K_{dp}^{1-way} (°/km) which is the range derivative of the differential propagation phase Ψ_{dp} . Furthermore, λ is the wavelength, $|K|^2$ is a dielectric factor, $\sigma_{hh/vv}$ are the radar cross sections, and $f_{hh/vv}$ are the forward scattering amplitudes of the raindrops at linear horizontal and linear vertical polarisations. For the purpose of this contribution, numerical values of the radar cross sections and the forward scattering amplitudes of raindrops have been computed for a frequency of 9.475 GHz (IDRA) by the Fredholm Integral Method, Holt et al. (1978), assuming dielectric spheroids with the axis ratio given by Illingworth et al. (2000), and the complex permittivity of water at a temperature of 10 °C, Liebe et al. (1991).

The propagation of electromagnetic waves through rain leads to attenuation of the measured reflectivity and the differential reflectivity, Fig. 1:

$$Z_{hh,att}(r_n) = Z_{hh}(r_n) - 2 \int_{r_1}^{r_n-1} \alpha_{hh}^{1-way}(r) dr \quad (6)$$

$$Z_{dr,att}(r_n) = Z_{dr}(r_n) - 2 \int_{r_1}^{r_n-1} \alpha_{h-v}^{1-way}(r) dr. \quad (7)$$

The specific attenuation α_{hh}^{1-way} (dB/km) and the differential attenuation α_{h-v}^{1-way} (dB/km) in the equations above are given by

$$\alpha_{hh}^{1-way} = 8.686 \lambda \int_D \Im[f_{hh}(D)] N(D) dD \quad (8)$$

and

$$\alpha_{h-v}^{1-way} = 8.686 \lambda \int_D \Im[f_{hh}(D) - f_{vv}(D)] N(D) dD. \quad (9)$$

A method to correct polarimetric weather radar measurements for attenuation has been introduced by Chandra et al. (1990). It is based on the estimation of the RDSD parameters that are subsequently used to determine the corresponding attenuations by scattering computations. The block diagram of this method, shown in Fig. 2, is explained next.

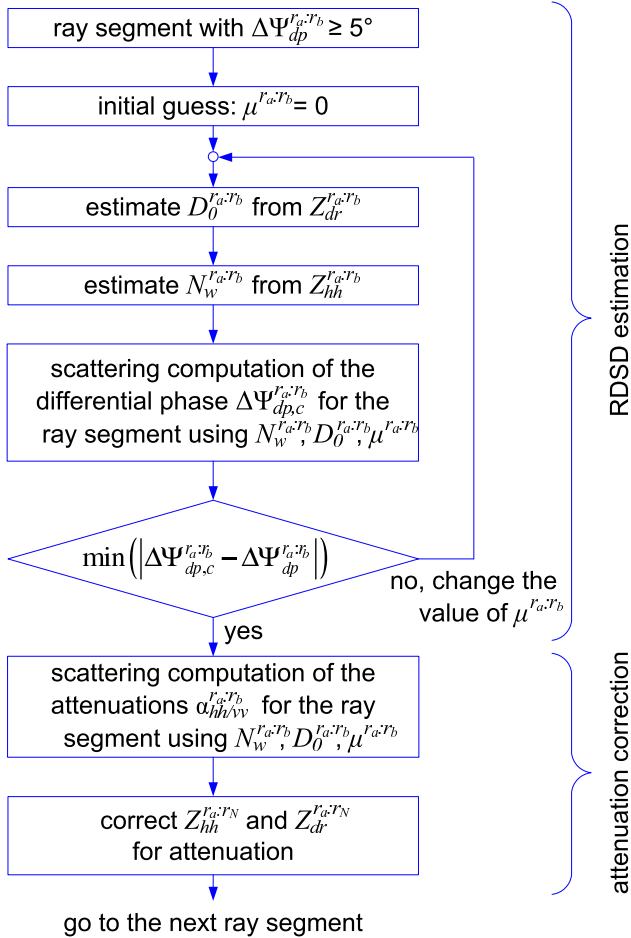


Fig. 2. Block diagram of the method of Chandra et al. (1990) to extract the RDSD parameters, and, subsequently, to correct for attenuation.

A measured ray of the integrated weather radar observables $Z_{hh,att}$, $Z_{dr,att}$ and Ψ_{dp} , such as the one exemplary shown in Fig. 1, is divided into ray segments satisfying $\Delta\Psi_{dp}^{r_a:r_b} = \Psi_{dp}(r_b) - \Psi_{dp}(r_a) \geq 5^\circ$. A precondition for the application of this method is a given raindrop shape and a temperature to calculate the complex permittivity of water. Starting at close range to the radar, i.e. $r_a = r_1$, Fig. 1, the median volume diameter $D_0^{r_a:r_b}$ is estimated via pre-calculated look-up tables from the measured differential reflectivity $Z_{dr,att}^{r_a:r_b}$, assuming an initial shape parameter $\mu^{r_a:r_b} = 0$ (blue curve in Fig. 3). After determining $N_w^{r_a:r_b}$ again via pre-calculated look-up tables (red curve in Fig. 3), scattering computations with the determined RDSD parameters are carried out in order to compute the expected differential propagation phase $\Delta\Psi_{dp,c}^{r_a:r_b}$ for the ray segment. The optimum shape parameter $\mu^{r_a:r_b}$ for the ray segment is found by minimising iteratively the difference among the measured and the computed differential propagation phase of the ray segment, i.e. $\min(|\Delta\Psi_{dp,c}^{r_a:r_b} - \Delta\Psi_{dp}^{r_a:r_b}|)$. The range of possible shape pa-

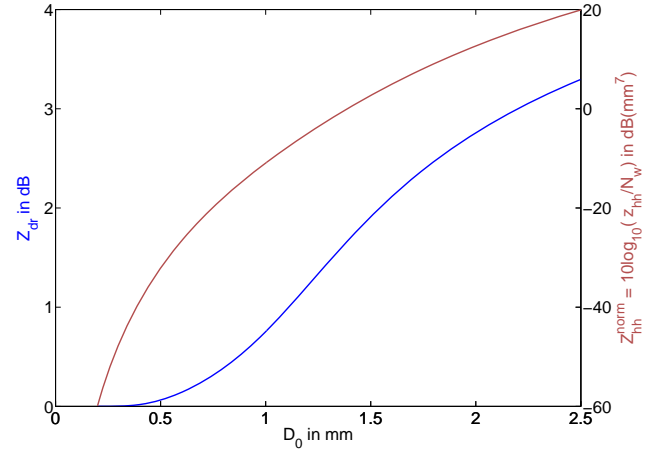


Fig. 3. Results of scattering computations to relate the differential reflectivity (blue curve) and the normalised reflectivity at horizontal polarisation (red curve) to the median volume diameter. The shape parameter is $\mu = 0$, and the raindrop axis ratio of Illingworth et al. (2000) was used.

rameters was limited within $-2 \leq \mu^{r_a:r_b} \leq 10$. Once the RDSD is characterised in this way, scattering computations are carried out again in order to estimate the attenuations, Eqs. (8) and (9), for the ray segment. After correcting $Z_{hh,att}^{r_a:r_N}$ and $Z_{dr,att}^{r_a:r_N}$ for the attenuation, the method is applied to the adjacent ray segment, i.e. $r_a = r_{b+1}$.

This method is only suited for rain, i.e. it is likely to fail in the presence of ground clutter or other hydrometeors such as graupel, hail or snow. Furthermore, Otto et al. (2009) showed that the application of this method is limited for IDRA to $0.5 \text{ dB} \leq Z_{dr} \leq 2.7 \text{ dB}$ in order to estimate the median volume diameter within $\pm 0.15 \text{ mm}$. Thus, the method is not suited for light rain (rain rates $< 3 \text{ mm/h}$) as well as for strong rain (rain rates $> 25 \text{ mm/h}$). In the case of strong rain, scattering computations show that the attenuation at X-band can be alternatively estimated by

$$\alpha_{hh} (\text{dB/km}) = 0.25 K_{dp} (^\circ/\text{km}) \quad (10)$$

and

$$\alpha_{h-v} (\text{dB/km}) = 0.05 K_{dp} (^\circ/\text{km}). \quad (11)$$

However, Eqs. (10) and (11) introduce uncertainties due to the variability of the microphysics of precipitation, Bringi et al. (1990), and due to the estimation of K_{dp} , Georgiou et al. (2009).

For each ray segment the RDSD is estimated from attenuated reflectivity and differential reflectivity measurements which means that the median volume diameter will be underestimated, blue curve in Fig. 3. This in turn results in an underestimation of Z_{hh}^{norm} , red curve in Fig. 3, and an overestimation of the concentration since $N_w = 10(Z_{hh,att} - Z_{hh}^{norm})/10 (\text{mm}^{-1} \text{ m}^{-3})$. At X-band this may yield for $\mu = 0$ to an underestimation of 0.3 mm for D_0 , and an

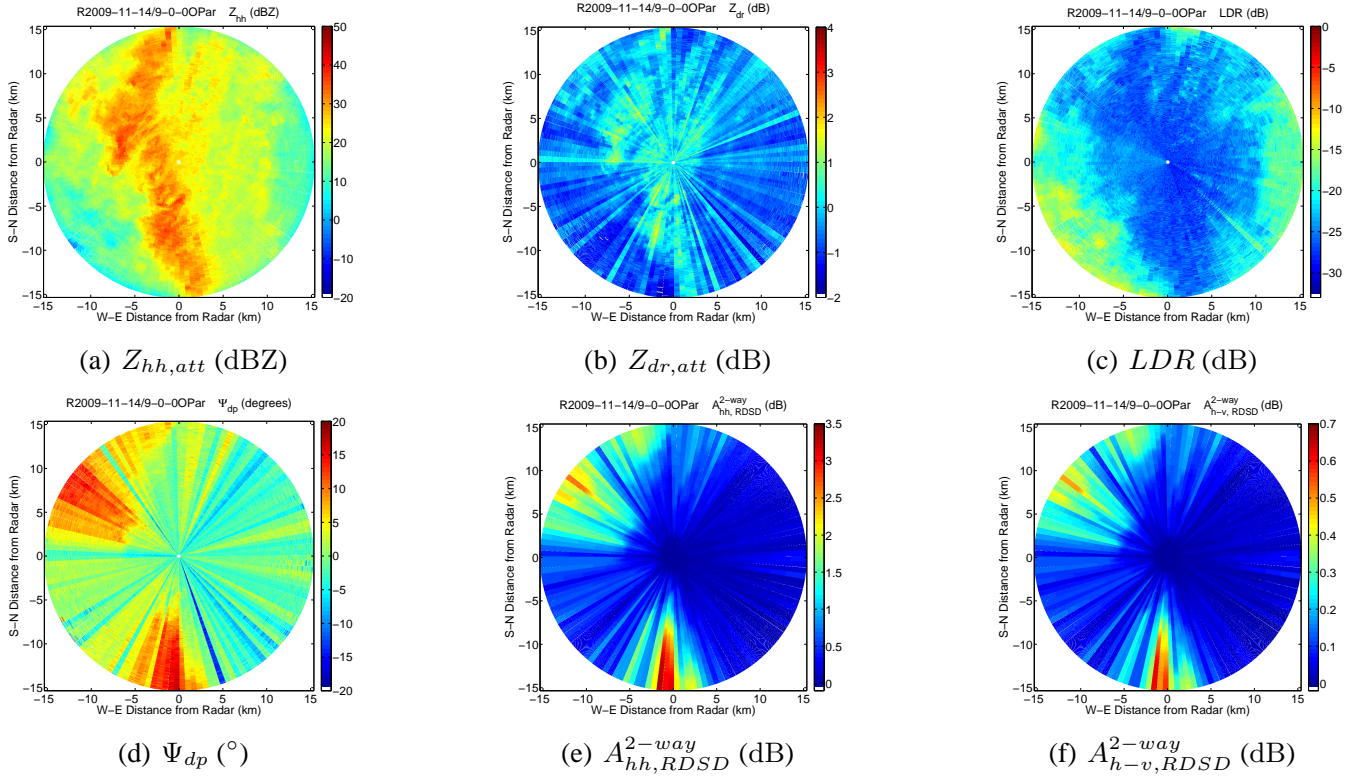


Fig. 4. IDRA precipitation measurements from the 14 November 2009 at 09:00 UTC. Shown are PPI's of (a) the reflectivity $Z_{hh,att}$ (dBZ), (b) the differential reflectivity $Z_{dr,att}$ (dB), (c) the linear depolarisation ratio LDR (dB) and (d) the differential propagation phase Ψ_{dp} (°). (e) and (f) show the 2-way cumulative attenuation at horizontal polarisation and the differential attenuation, respectively, derived by the method of Chandra et al. (1990).

overestimation of N_w up to a factor of 100. For the computation of the attenuations, Eqs. (8) and (9), this may yield in turn to an over- or underestimation. In any case, the RDSD parameters are estimated again after the first run of this method from the attenuation corrected reflectivities. The presented method is applied in the next section to IDRA measurements in order to assess the accuracy of the estimated attenuations. Therefore, they are compared to the attenuations derived by the phase constraint method with imposed self-consistency of Gorgucci and Baldini (2007).

3 Application to IDRA data

Figure 4 shows a precipitation measurement of IDRA from the 14 November 2009, 09:00 UTC. IDRA is located in the centre of these plan-position indicators (PPI), that from Fig. 4a to d illustrate the measured weather radar observables reflectivity $Z_{hh,att}$ (dBZ), differential reflectivity $Z_{dr,att}$ (dB), linear depolarisation ratio LDR (dB), and differential propagation phase Ψ_{dp} (°). The presence of propagation effects is not evident in the reflectivity and the differential reflectivity measurement. However, the measured differential phase,

Fig. 4d, shows cumulative values of up to 20° over the range of 15 km in some sectors.

Using the measured $Z_{hh,att}$, $Z_{dr,att}$ and Ψ_{dp} , the attenuations were determined by the method of Chandra et al. (1990) introduced in the preceding section. The resulting cumulative attenuation at horizontal polarisation and the differential attenuation are shown in Fig. 4e and f.

For a comparison, the attenuations were also derived by the method outlined in Gorgucci and Baldini (2007). This method relies on a first estimation of the attenuations by the application of the phase constraint method of Testud et al. (2000). Then the self-consistency equations

$$\alpha_{hh} = a_1 z_{hh}^{b_1} z_{dr}^{c_1} K_{dp}^{d_1} \quad (12)$$

and

$$\alpha_{h-v} = a_2 z_{hh}^{b_2} z_{dr}^{c_2} K_{dp}^{d_2} \quad (13)$$

are imposed as outlined in Gorgucci and Baldini (2007), resulting in an adjustment of the estimated attenuations. The reflectivities z_{hh} and z_{dr} in Eqs. (12) and (13) are in linear units, and the coefficients of these self-consistency equations for the frequency of IDRA can be found in Table 1. In the following comparison of the attenuations, the subscript SCPC

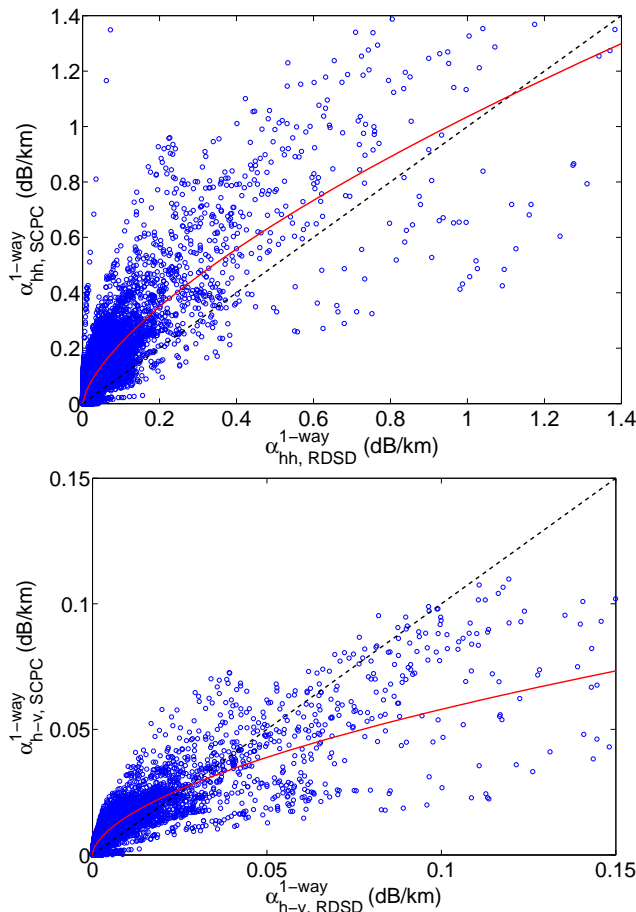


Fig. 5. Scatter plots comparing (a) the specific attenuations at horizontal polarisation $\alpha_{hh}^{1\text{-way}}$, and (b) the differential attenuations $\alpha_{h-v}^{1\text{-way}}$ estimated by the methods of Chandra et al. (1990) (RDSD) and Gorgucci and Baldini (2007) (SCPC) for 35317 range bins of two PPI's measured by IDRA. The dashed black line corresponds to $x = y$. The red line is a power-law fit.

denotes attenuations derived by the self-consistency method of Gorgucci and Baldini (2007), and the subscript RDSD denotes attenuations derived by the method of Chandra et al. (1990).

Figure 5 shows two scatter plots comparing the attenuations $\alpha_{hh}^{1\text{-way}}$ and $\alpha_{h-v}^{1\text{-way}}$ estimated by the two methods presented above from IDRA data. The data comprise 35317 range bins from two PPI's measured the 25 May 2009, 20:23 UTC, and the 14 November 2009, 09:00 UTC, Fig. 4. The red lines in the scatter plots are power law fits $y = mx^n$ to the data with $m_{hh} = 1$, $n_{hh} = 0.67$, and $m_{h-v} = 0.22$, $n_{h-v} = 0.58$. The general tendency shows that for small attenuations $\alpha_{hh}^{1\text{-way}} < 1.1$ dB/km and $\alpha_{h-v}^{1\text{-way}} < 0.03$ dB/km, the RDSD method is underestimating the attenuations compared to SCPC method, whereas larger attenuations are overestimated. This non-linear correlation between the estimated attenuations that is most notably in Fig. 5b may be attributed to the uncertainties introduced by estimating the RDSD pa-

Table 1. Coefficients for the self-consistency Eqs. (12) and (13) derived for a frequency of 9.475 GHz for a set of 1500 normalised raindrop-size distributions with varying N_w , D_0 , μ , raindrop axis ratios and temperatures.

a_1	b_1	c_1	d_1
4.95×10^{-3}	0.463	-1.321	0.539
a_2	b_2	c_2	d_2
7.58×10^{-4}	0.378	0.416	0.64

rameters from attenuated reflectivities. Depending on the actual values of the reflectivities, the resulting opposite errors of D_0 and N_w do show non-linearities, Fig. 3. One may think of additional iterations within the RDSD method to improve the estimated attenuations. However, this would yield an increase of the processing time. Instead, at X-band it is suggested to perform the attenuation correction prior to the estimation of the RDSD parameters.

4 Conclusions

The purpose of this contribution was the evaluation of a method to estimate the RDSD parameters and in the same vein to perform a correction for attenuation for polarimetric weather radar data at X-band, Chandra et al. (1990). The focus was laid on the estimated attenuations by comparing them to a phase-constraint attenuation correction method with imposed self-consistency, Gorgucci and Baldini (2007). Small attenuations were underestimated by the RDSD method compared to the SCPC method whereas large attenuations were overestimated by the RDSD method. The attenuations estimated by the RDSD method result from scattering computations with RDSD parameters that are estimated from attenuated reflectivities. This may result in an underestimation of the median volume diameter D_0 but an overestimation of the concentration N_w that may explain the observed deviations of the estimated attenuations. Since the RDSD method is only applicable for a limited differential reflectivity range, no attenuation estimate is provided for weak and for strong rain. For X-band, this suggests to apply the attenuation correction and the RDSD estimation separately, i.e. first correct the reflectivities for attenuation using a phase-constraint method, followed by the estimation of the RDSD parameters. However it must be stressed, that the results of the comparison of the attenuation correction methods in this contribution can not be regarded as absolute. As a next step, we shall take advantage of the operational C-band weather radar of the Royal Netherlands Meteorological Institute at De Bilt that is situated in the vicinity of IDRA in a distance of approximately 25 km. The analysis and comparison of simultaneous measurements of precipitation events shall enable us to estimate the performance of attenuation correction methods.

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